

1 On the impact of wind farms on a convective atmospheric 2 boundary layer

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5 **Abstract.** With the rapid growth in the number of wind turbines installed world-
6 wide, a demand exists for a clear understanding of how wind farms modify land-
7 atmosphere exchanges. Here, we conduct three-dimensional large-eddy simulations
8 to investigate the impact of wind farms on a convective atmospheric boundary layer.
9 Surface temperature and heat flux are determined using a surface thermal energy
10 balance approach, coupled with the solution of a three-dimensional heat equation in
11 the soil. We study several cases of aligned and staggered wind farms with different
12 streamwise and spanwise spacings. The farms consist of Siemens SWT-2.3-93 wind
13 turbines. Results reveal that, in the presence of wind turbines, the stability of the
14 atmospheric boundary layer is modified, the boundary layer height is increased, and
15 the magnitude of the surface heat flux is slightly reduced. Results also show an
16 increase in land-surface temperature, a slight reduction in the vertically-integrated
17 temperature, and a heterogeneous spatial distribution of the surface heat flux.

18 **Keywords:** Convective atmospheric boundary layer, Large-eddy simulation, Wind
19 farm

20 1. Introduction

21 The wind field in the lowest part of the atmosphere is the most im-
22 portant atmospheric factor for wind-energy applications. A number of
23 recent studies (e.g., Baidya Roy et al., 2004; Calaf et al., 2011; Porté-
24 Agel et al., 2011; Lu and Porté-Agel, 2011; Fitch et al., 2012, 2013;
25 Abkar and Porté-Agel, 2013) have examined the interaction between
26 atmospheric boundary-layer (ABL) flow and wind farms, and generally
27 found that wind-turbine blade motions reduce wind speed, enhance
28 turbulence, and change the stability of the ABL flow. Warming effects,
29 particularly at nighttime, have been reported in a large-eddy simula-
30 tion (LES) study of a wind farm in a stably-stratified ABL (Lu and
31 Porté-Agel, 2011). Using satellite data, Zhou et al. (2012) have found
32 a significant warming trend of up to 0.7 °C on the land surface. Some
33 mesoscale simulations (e.g., Baidya Roy et al., 2004; Baidya Roy, 2011)

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34 have indicated cooling by wind farms during daytime; however, a high-
35 resolution study of wind-farm impacts on the convective atmospheric
36 boundary layer (CBL) has not been conducted to date. Considering
37 the fast worldwide expansion of wind energy, understanding the interac-
38 tions between wind farms and the ABL is important for predicting their
39 performance, quantifying their impacts on local meteorology, improving
40 their parametrisation in weather models, and assessing their effects on
41 collocated agricultural crops (e.g., due to local changes in temperature,
42 evaporation and transpiration).

43 Computational fluid dynamics simulations can be used to study
44 complex engineering and environmental turbulent flows, where con-
45 trolled measurements are difficult or impossible to perform, especially
46 for very large systems such as wind farms. Although several mesoscale
47 simulations (Baidya Roy et al., 2004; Baidya Roy, 2011; Fitch et al.,
48 2012, 2013) have been performed to estimate large-scale impacts of
49 wind farms, they do not provide insight into the flow details near the
50 land surface, where effects on turbulent fluxes are important. Also, most
51 simulations (e.g., Baidya Roy et al., 2004; Calaf et al., 2011; Baidya
52 Roy, 2011; Fitch et al., 2012, 2013) do not consider the wind-turbine-
53 induced rotation forces, and assume uniform force distribution over
54 the rotor plane. Wake rotation plays an important role in wind-turbine
55 mixing (Lu and Porté-Agel, 2011; Porté-Agel et al., 2011; Markfort
56 et al., 2012; Zhang et al., 2013a). Failure to take into account wake-
57 rotation effects has been shown (Porté-Agel et al., 2011; Zhang et al.,
58 2013a) to result in errors in the prediction of momentum and heat
59 fluxes near the land surface. Parametrisation of wind turbines using
60 the actuator line model (ALM) and actuator disk with rotation model
61 is capable of reproducing important turbulent wake features, such as
62 the formation of helicoidal tip vortices (with the ALM), the enhanced
63 turbulence levels at the top edge of the wakes, and the rotation of the
64 wakes (Lu and Porté-Agel, 2011; Porté-Agel et al., 2011; Wu and Porté-
65 Agel, 2011). Specifically, Lu and Porté-Agel (2011) used LES with the
66 ALM to investigate the effects of a large wind farm on a stably-stratified
67 ABL.

68 In this study, we investigate, for the first time, a dry CBL flow
69 through extensive wind farms, with emphasis on the characteristics
70 of wind-turbine wakes and their aggregated effect on land-atmosphere
71 exchange (momentum and heat fluxes). The LES framework and the
72 simulation details are described in Sect. 2, and results are presented
73 and discussed in Sects. 3-5. A summary and conclusions are provided
74 in Sect. 6.

2. Large-eddy simulation framework

75

2.1. LES GOVERNING EQUATIONS AND MODELS

76

77 We use a modified LES code that has been used in previous wind-
 78 energy studies (Porté-Agel et al., 2011; Wu and Porté-Agel, 2011; Lu
 79 and Porté-Agel, 2014). We aim to study the dynamics of a dry CBL
 80 that excludes moisture, and solve the filtered continuity equation, the
 81 filtered momentum conservation equations based on the Boussinesq
 82 approximation, and the filtered transport equation for potential tem-
 83 perature,

$$84 \quad \frac{\partial \tilde{u}_i}{\partial x_i} = 0, \quad (1)$$

85

$$86 \quad \frac{\partial \tilde{u}_i}{\partial t} + \tilde{u}_j \frac{\partial \tilde{u}_i}{\partial x_j} = -\frac{\partial \tilde{p}}{\partial x_i} - \frac{\partial \tau_{ij}}{\partial x_j} + \frac{\tilde{\theta} - \langle \tilde{\theta} \rangle_h}{\Theta_0} \delta_{i3} g + f_c \varepsilon_{ij3} (\tilde{u}_j - U_{G,j}) + \mathcal{F}_i, \quad (2)$$

87

$$88 \quad \frac{\partial \tilde{\theta}}{\partial t} + \tilde{u}_j \frac{\partial \tilde{\theta}}{\partial x_j} = -\frac{\partial q_j}{\partial x_j}, \quad (3)$$

89 where the tilde ($\tilde{\cdot}$) represents a spatial filtering at the resolved scale
 90 $\tilde{\Delta}$, $(\tilde{u}_1, \tilde{u}_2, \tilde{u}_3) = (\tilde{u}, \tilde{v}, \tilde{w})$ are the components of the velocity field, $\tilde{\theta}$ is
 91 the resolved potential temperature, $\tau_{ij} = \tilde{u}_i \tilde{u}_j - \tilde{u}_i \tilde{u}_j$ is the subgrid-
 92 scale (SGS) stress tensor, $q_i = \tilde{u}_i \tilde{\theta} - \tilde{u}_i \tilde{\theta}$ is the SGS flux vector, Θ_0
 93 is the reference temperature, $\langle \cdot \rangle_h$ represents a horizontal average, g
 94 is the acceleration due to gravity, f_c is the Coriolis parameter, $U_{G,i}$
 95 is the geostrophic wind speed, δ_{ij} is the Kronecker delta, ε_{ijk} is the
 96 alternating unit tensor, \tilde{p} is the effective pressure, and \mathcal{F}_i is a forcing
 97 term (e.g., wind-turbine induced forces). Since only dry air conditions
 98 are simulated (water vapour and cloud formation are not included),
 99 radiative heating and cooling in the air can be neglected (Arya, 2001;
 100 Holton, 2004).

101 We adopt the dynamic version of the recently-developed modulated
 102 gradient models (Lu and Porté-Agel, 2010, 2013, 2014) for the SGS
 103 stress and for the SGS flux vector. The turbine-induced forces are
 104 parametrised using the rotating actuator disk model, which accounts
 105 for the effect of the turbine-induced flow rotation as well as the non-
 106 uniform force distribution (Wu and Porté-Agel, 2011). Figure 1 shows
 107 a cross-section airfoil element in the (θ, x) plane, where x is the axial
 108 direction. Denoting the tangential and axial velocity in the inertial
 109 frame of reference as V_θ and V_x , respectively, the local velocity relative
 110 to the rotating blade is given as $\mathbf{V}_{rel} = (V_\theta - \Omega r, V_x)$. The angle of
 111 attack is defined as $\alpha = \varphi - \gamma$, where $\varphi = \tan^{-1}(V_x/(\Omega r - V_\theta))$ is the

112 angle between V_{rel} and the rotor plane, and γ is the local pitch angle.
 113 The resulting force is given by

$$114 \quad \mathbf{f}_{2D} = \frac{d\mathbf{F}}{dA} = \frac{1}{2}\rho V_{rel}^2 \frac{Bc}{2\pi r} (C_L \mathbf{e}_L + C_D \mathbf{e}_D), \quad (4)$$

115 where an annular area of differential size is $dA = 2\pi r dr$, r is the radius,
 116 V_{rel} is the local velocity relative to the rotating blade, B is the number
 117 of blades, $C_L = C_L(\alpha, Re)$ and $C_D = C_D(\alpha, Re)$ are the lift coefficient
 118 and the drag coefficient, respectively, c is the chord length, and \mathbf{e}_L and
 119 \mathbf{e}_D denote the unit vectors in the directions of the lift and the drag,
 120 respectively.

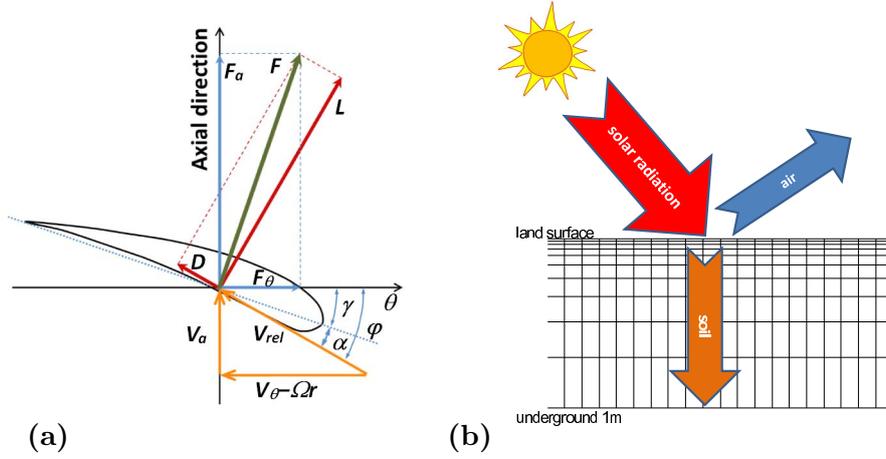


Figure 1. (a) Cross-section airfoil element showing velocities and force vectors; (b) schematic of energy balance, and structured grid with logarithmic vertical spacing used to solve the heat equation in the soil up to a depth of 1 m.

121 At the surface, the instantaneous wall stress is related to the velocity
 122 at the first vertical node through the application of the Monin-Obukhov
 123 similarity theory (Businger et al., 1971; Stull, 1988; Arya, 2001). Al-
 124 though this typically applies to mean quantities, it is common practice
 125 (Lu and Porté-Agel, 2010) in LES of atmospheric flows to use it for
 126 instantaneous fields as follows,

$$127 \quad \tau_{i3}|_w = -u_*^2 \frac{\tilde{u}_i}{u_r} = - \left(\frac{u_r \kappa}{\ln(z/z_0) - \Psi_M} \right)^2 \frac{\tilde{u}_i}{u_r} \quad (i = 1, 2), \quad (5)$$

128 where $\tau_{i3}|_w$ is the instantaneous local wall stress, u_* is the friction ve-
 129 locity, z_0 is the aerodynamic roughness, κ is the von Kármán constant,
 130 Ψ_M is the stability correction for momentum, and u_r is the local filtered
 131 horizontal velocity at the first vertical level. In a similar manner, the

132 surface heat flux is computed as

$$133 \quad q_3|_w = u_* \theta_* = \frac{u_* \kappa (\theta_s - \tilde{\theta})}{\ln(z/z_{0,\theta}) - \Psi_H}, \quad (6)$$

134 where $z_{0,\theta}$ is the aerodynamic roughness for the potential temperature,
 135 θ_* is a temperature scale, and θ_s is the surface (ground level) potential
 136 temperature. Following Stull (1988), Arya (2001), we adopt $\Psi_M =$
 137 $2 \ln\left(\frac{1+x}{2}\right) + \ln\left(\frac{1+x^2}{2}\right) - 2 \tan^{-1}(x) + \frac{\pi}{2}$ and $\Psi_H = 2 \ln\left(\frac{1+x^2}{2}\right)$, where
 138 $x = \left(1 - \frac{15z}{L}\right)^{1/4}$, and $L = -\frac{u_*^3 \tilde{\theta}}{\kappa g q_3|_w}$ is the Obukhov length. As adopted
 139 in previous studies (e.g., Kosovic and Curry, 2000; Beare et al., 2006),
 140 the boundary-layer height, z_i , is computed as (1/0.95) times the height
 141 at which the horizontally-averaged stress falls to 5% of its surface value.
 142 As in previous studies (e.g., Beare et al., 2006), a Rayleigh damping
 143 layer is set above 1200 m to limit gravity-wave reflection from the top
 144 of the domain.

145 To determine land-surface temperature and surface heat flux, we
 146 adopt a surface thermal energy balance approach along with 10 levels
 147 of soil temperature to a depth of 1 m using a logarithmic spacing, as
 148 shown in Fig. 1b. The surface thermal energy balance can be written
 149 as

$$150 \quad R_N = \frac{c_s \alpha}{c_a} \frac{\partial \tilde{\theta}}{\partial z} \Big|_w + q_3|_w, \quad (7)$$

151 where R_N is the net solar radiation, c_s and c_a are the heat capacities
 152 of the soil and the air, and α is the diffusivity coefficient of the soil. A
 153 similar method was used in Deardorff (1974), who assumed horizontal
 154 homogeneity and solved the one-dimensional heat equation in the soil.
 155 In this study we relax the assumption of horizontal homogeneity, and
 156 solve a three-dimensional heat equation in the soil

$$157 \quad \frac{\partial \tilde{\theta}}{\partial t} = \alpha \nabla^2 \tilde{\theta}, \quad (8)$$

158 which allows us to capture the heterogeneity of the land-surface tem-
 159 perature and the surface heat flux.

160 2.2. NUMERICAL SET-UP

161 In order to understand the impact of a wind farm on a CBL flow,
 162 we first simulate a baseline CBL case (without wind turbines). We
 163 have revised the numerical procedures adopted for other ABL flow
 164 cases (Moeng, 1984; Mason, 1989; Moeng and Sullivan, 1994; Agee

165 and Gluhovsky, 1999; Sorbjan, 2006; Beare et al., 2006; Conzemius
 166 and Fedorovich, 2006, 2008) to make the simulated CBL suitable for
 167 studying its interactions with wind farms. In summary, the bound-
 168 ary layer is driven by an imposed uniform geostrophic wind speed of
 169 15 m s^{-1} ; the Coriolis parameter is set to $f_c = 1.00 \times 10^{-4} \text{ rad s}^{-1}$,
 170 corresponding to a latitude of about 45° ; $z_0 = 0.1 \text{ m}$ and $z_{0,\theta} = 0.01 \text{ m}$;
 171 $g = 9.81 \text{ m s}^{-2}$; $\theta_0 = 300 \text{ K}$. The initial potential temperature profile
 172 consists of a mixed layer (with potential temperature 302 K , which is
 173 also the initial soil temperature) up to 100 m with an overlying inversion
 174 of strength 0.0114 K m^{-1} . We assume the net solar radiation has a con-
 175 stant value of 0.08 K m s^{-1} during the simulation. The soil is dry, and
 176 its diffusivity is $5.0 \times 10^{-7} \text{ m}^2 \text{ s}^{-1}$ (Deardorff, 1974; Stull, 1988). The
 177 heat capacity of the soil is $1.3 \times 10^6 \text{ J m}^{-3} \text{ K}^{-1}$ (Stull, 1988). It should
 178 be noted that the boundary-layer height of the baseline CBL case is
 179 continuously increasing. According to previous time-scale arguments
 180 (Agee and Gluhovsky, 1999), the baseline CBL case is fully developed
 181 after 3 h (approximately 18 large-eddy turnover times). Therefore,
 182 in order to examine the wind-turbine effects relative to the baseline
 183 case, we introduce the wind turbines only after 3 h , and the mean
 184 velocity direction is aligned to be axial at the hub-height level. Note
 185 that during the wind-farm simulation, the wind-direction change in the
 186 wind-turbine region is not significant.

187 The domain is uniformly divided into N_x , N_y and N_z grid points
 188 in the x , y and z directions. Periodic boundary conditions are applied
 189 horizontally so that an idealized very large (effectively infinite) wind
 190 farm can be simulated. A pseudo-spectral method is adopted in the
 191 horizontal directions, and vertical derivatives are approximated with
 192 second-order central differences. The vertical domain has a height of
 193 $L_z = 1476 \text{ m}$, and the vertical grid number is $N_z = 128$. The grid
 194 planes are staggered in the vertical with the first vertical velocity plane
 195 at a distance $\Delta z = \frac{L_z}{N_z - 1}$ from the surface. Aliasing errors are corrected
 196 in the nonlinear terms using the $\frac{3}{2}$ rule (Canuto et al., 1988). The
 197 time advancement is carried out using a second-order accurate Adams-
 198 Bashforth scheme (Canuto et al., 1988). We set a constant timestep
 199 corresponding to a rather restrictive Courant-Friedrichs-Lewy number
 200 of about 0.02 to reduce the error from the time stepping.

201 Siemens SWT-2.3-93 wind turbines, with a rotor diameter (D) of
 202 93 m and a hub height of 80 m , are ‘immersed’ in the flow. Details
 203 of the wind turbine can be found in Leloudas (2006), and Laursen
 204 et al. (2007). According to previous domain-size arguments (Roode
 205 and Duynkerke, 2004), a horizontal domain size of $1 \sim 2$ times of
 206 the boundary-layer height is sufficient for a dry CBL simulation. In
 207 this study, the horizontal domain is approximately four times of the

208 boundary-layer height. We vary the horizontal dimensions (L_x and L_y),
 209 the resolutions (N_x and N_y), the number of wind turbines ($N_{t,x}$ by
 210 $N_{t,y}$), the layout (aligned or staggered), and the distance between wind
 211 turbines ($S_x D$ by $S_y D$). The suite of LES cases is described in Table
 212 I. For simplicity, the aligned $S_x \times S_y = X \times Y$ wind farm case is
 213 abbreviated to ‘aX×Y,’ and the staggered $S_x \times S_y = X \times Y$ wind
 214 farm case is abbreviated to ‘sX×Y.’ As an example, Figs. 2 and 3
 215 show instantaneous fields and wind-turbine induced vortices in two
 wind-farm cases.

Table I. Parameters of the wind farm cases.

$S_x \times S_y$	$N_{t,x} \times N_{t,y}$	$L_x \times L_y$ [m^2]	$N_x \times N_y$
7×7	6×5	3906×3255	168×280
6×6	7×6	3906×3348	168×288
5×5	8×7	3720×3255	160×280

216

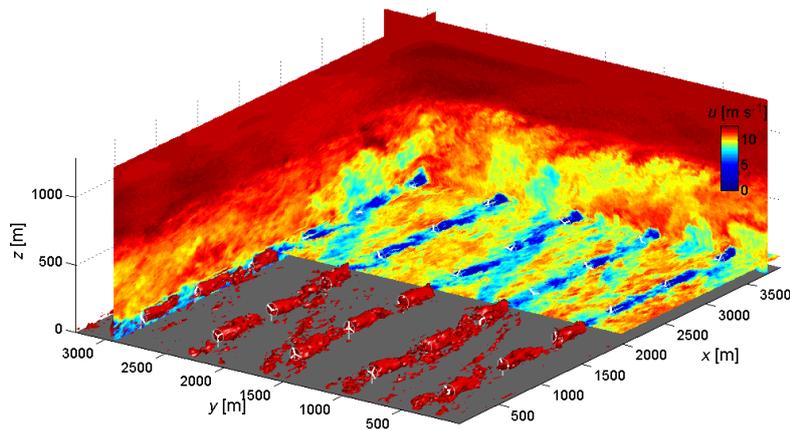


Figure 2. Flow field in a fully developed wind-turbine array, a7×7 case, shown by streamwise velocity contours (plotted on three representative (x, z) -, (y, z) -, and (x, y) -planes) and iso-surface of vorticity.

217

3. Mean vertical profiles

218 Figure 4 shows vertical profiles of the mean wind speed (defined as
 219 $\left(\langle \tilde{u} \rangle^2 + \langle \tilde{v} \rangle^2\right)^{\frac{1}{2}}$, where the angle brackets, $\langle \cdot \rangle$, represent averaging over

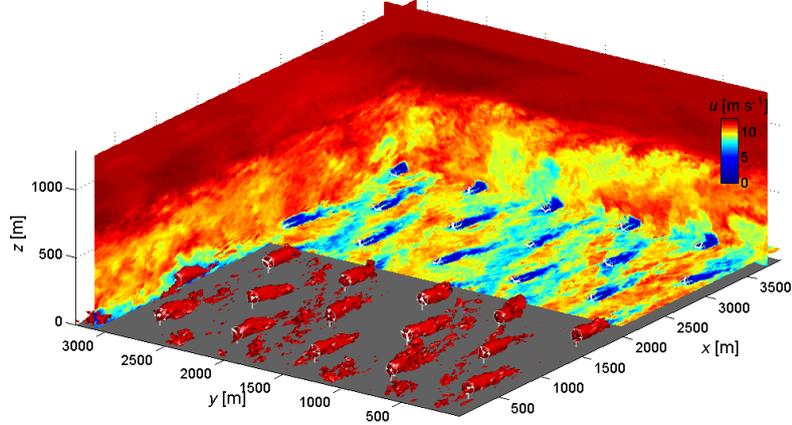


Figure 3. Flow field in a fully developed wind-turbine array, $s7 \times 7$ case, shown by streamwise velocity contours (plotted on three representative (x, z) -, (y, z) -, and (x, y) -planes) and iso-surface of vorticity.

220 1 hr and horizontal directions) and the mean potential temperature
 221 obtained from the $a5 \times 5$ wind-farm case and the baseline (no-farm) case.
 222 Results clearly reveal the extraction of kinetic energy by the turbines.
 223 The presence of the wind farm increases the boundary-layer height
 224 by approximately 150 m (about 16%) after 10 h. It also leads to an
 225 increase of about 0.5 K in land-surface temperature, and a slight re-
 226 duction of about 0.03 K in the vertically-integrated CBL temperature,
 227 which is consistent with the reduction in the surface heat flux (shown
 228 later). Table II presents the final (at 10 h) changes of land-surface
 229 and vertically-integrated temperatures induced by the wind farm for
 230 all the layouts considered here. The table shows that denser wind-farm
 231 layouts bear larger differences with respect to the baseline (no-farm)
 232 case. Moreover, for a given turbine density, the staggered wind farm
 233 bears larger difference than its aligned counterpart. The primary reason
 234 is that, compared to the aligned counterpart, the staggered wind-farm
 235 configuration yields more energy extraction (shown later). This yields
 236 greater downward momentum transport and more efficient mixing, as
 237 shown in recent studies (e.g., Markfort et al., 2012).

238 Figure 5 compares vertical profiles of the total turbulent momentum
 239 flux (defined as $\left(\langle \tilde{u}'\tilde{w}' + \tau_{13} \rangle^2 + \langle \tilde{v}'\tilde{w}' + \tau_{23} \rangle^2\right)^{\frac{1}{2}}$, where the resolved
 240 fluctuation of an arbitrary variable, a , is written as $\tilde{a}' = \tilde{a} - \langle \tilde{a} \rangle$) and
 241 the total turbulent heat flux (defined as $\langle \tilde{\theta}'\tilde{w}' + q_3 \rangle$) obtained from the

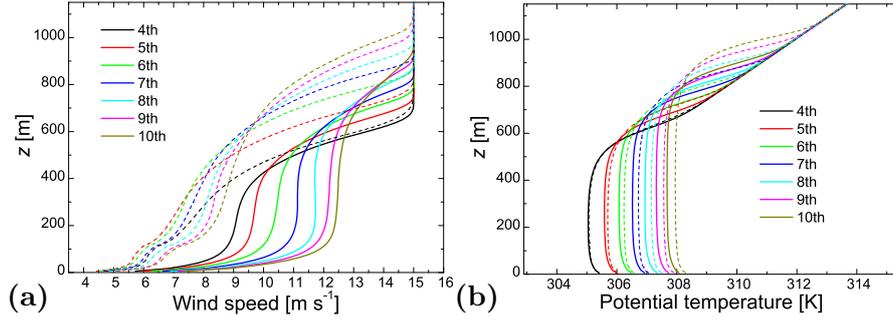


Figure 4. Vertical profiles of (a) mean wind speed and (b) mean potential temperature, obtained from the baseline case (solid line) and the a5×5 case (dashed line).

Table II. Final (10 h) temperature difference between wind-farm and baseline (no-farm) cases. $\Delta\theta_S$: temperature change on the land surface; $\Delta\theta_{BL}$: temperature change over the boundary layer (vertically-integrated mean over 0-1200 m).

Case	s5×5	a5×5	s6×6	a6×6	s7×7	a7×7
$\Delta\theta_S$ [K]	0.650	0.623	0.524	0.512	0.429	0.414
$\Delta\theta_{BL}$ [K]	-0.0365	-0.0336	-0.0289	-0.0269	-0.0236	-0.0215

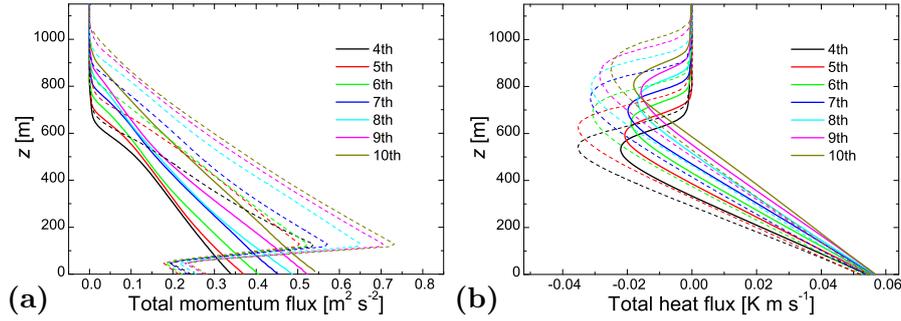


Figure 5. Vertical profiles of (a) total momentum flux and (b) total heat flux, obtained from the baseline case (solid line) and the a5×5 case (dashed line).

242 a5×5 wind-farm case and the baseline case. In the baseline case, the
 243 momentum flux shows a near-linear decrease in magnitude with height.
 244 It is evident that the presence of the wind farm dramatically changes
 245 the momentum and heat flux profiles. From both profiles, it is also clear
 246 that the wind farm increases the boundary-layer height by about 16%.
 247 The surface momentum flux (and thus the friction velocity) is reduced
 248 due to the extraction of momentum by the wind turbines. In line with
 249 results from previous studies (e.g., Calaf et al., 2011; Markfort et al.,

250 2012), the maximum magnitude of the turbulent vertical momentum
251 flux is found at the turbine top-tip height. At that level, the high shear
252 found at the upper edge of the turbine wakes leads to high production
253 of turbulence kinetic energy ($\text{TKE} = (\sigma_u^2 + \sigma_v^2 + \sigma_w^2) / 2$, where σ_u^2 , σ_v^2
254 and σ_w^2 are the variances of the three velocity components) and, in
255 turn, large TKE and momentum flux.

256 The warming produced by wind farms under stable conditions is
257 caused by the enhanced vertical entrainment of relatively warmer air
258 from higher altitudes (Baidya Roy and Traiteur, 2010; Lu and Porté-
259 Agel, 2011; Zhou et al., 2012). Under convective conditions, current
260 results show that the land-surface temperature and the near-surface
261 temperature are increased, but the vertically-integrated temperature
262 is slightly reduced. In comparison with the ABL flows under stable
263 conditions, the mixing under convective conditions is already very large
264 even without turbines; hence, the turbine-enhanced turbulent mixing
265 plays a relatively smaller role in the temperature distribution.

266 The vertical profile of the turbulent heat flux, shown in Fig. 5b,
267 provides a better understanding of the thermal exchanges and conse-
268 quent temperature changes induced by the wind farms. In the current
269 scenario, wind-turbine blade motions lead to relatively smaller changes
270 in the turbulent heat flux near the surface, compared with the reduction
271 in heat flux magnitude previously reported under stable conditions (see
272 Fig. 19b in Lu and Porté-Agel, 2011). However, near the boundary-
273 layer top, the turbulent heat flux profiles reveal a largely enhanced
274 entrainment flux in the presence of wind turbines. Specifically, the
275 entrainment-flux to surface-flux ratio increases from 0.29 to 0.48 due
276 to the wind-farm effect. The enhanced entrainment at the top of the
277 boundary layer due to the presence of wind farms indicates increased
278 downward flux of relatively warmer air in the entrainment layer. As a
279 result, even though the near-surface temperature is slightly higher, the
280 temperature in the entrainment layer is considerably lower than that in
281 the baseline case. This is consistent with the fact that the wind turbines
282 act as large roughness elements producing a substantial enhancement of
283 momentum flux and turbulence kinetic energy at the wind-turbine top-
284 tip level and throughout the boundary layer. Moreover, the momentum
285 transfer to the wind-turbine region in very large wind farms is achieved
286 mainly by entraining warmer air from the free atmosphere. Also, an
287 increased shear at the entrainment layer is expected to lead to a larger
288 fraction of entrainment flux (relative to the surface flux) as shown
289 in previous studies of CBLs (e.g., Pino et al., 2003; Conzemius and
290 Fedorovich, 2006). This entrainment warming effect is compensated by
291 the reduced surface heat flux. This is in contrast to the stable boundary

292 layer case, for which both entrainment and surface fluxes contribute to
 293 the warming of the boundary layer.

294 Regarding the overall thermal-energy budget, the 4%-7% reduction
 295 in surface heat flux induced by the wind farm is consistent with the
 296 decrease in the vertically-integrated temperature. It also leads to larger
 297 heat flux into the soil (shown later). For the same turbine density, the
 298 differences between the surface fluxes for the staggered and aligned lay-
 299 outs are not significant. In contrast, the effect of wind-turbine density
 300 on the surface fluxes is evident. In particular, denser wind-farm layouts
 301 bear lower surface heat flux, which yields larger temperature differences
 302 as shown in Table II.

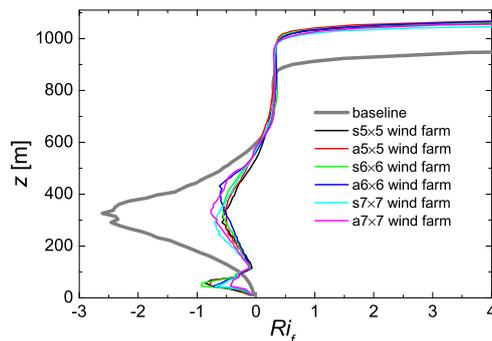


Figure 6. Vertical profiles of the flux Richardson number obtained from the baseline case and wind farm cases.

303 The Richardson number is an important dimensionless stability pa-
 304 rameter. In Fig. 6, we investigate stability changes by presenting the
 305 vertical profiles of the flux Richardson number,

$$306 \quad Ri_f = \frac{\frac{g}{\Theta_0} \langle \tilde{\theta}' \tilde{w}' + q_3 \rangle}{\langle \tilde{u}' \tilde{w}' + \tau_{13} \rangle \frac{\partial \langle \tilde{u} \rangle}{\partial z} + \langle \tilde{v}' \tilde{w}' + \tau_{23} \rangle \frac{\partial \langle \tilde{v} \rangle}{\partial z}} . \quad (9)$$

307 Results show that all wind-farm cases yield similar profiles of Ri_f .
 308 Below the wind turbine level, the flow is more unstable compared with
 309 the baseline case. It can be explained by the fact that the reduction in
 310 the magnitude of the momentum flux, due to the slow down of the flow
 311 induced by the wind-turbine blade motions in that region, is relatively
 312 stronger compared with the reduction in the heat flux. In the wind-
 313 turbine wake region and above, the flow is, in contrast, less unstable.
 314 The change of stability of the flow is associated with the turbulence
 315 enhancement produced by the wake flows. The wind-farm wake induces
 316 a higher momentum flux and a higher mean shear around the top-tip
 317 level, compared with the no-farm baseline case. This, together with

318 the reduction in the turbulent heat flux, makes the magnitude of the
 319 Richardson number (and the relative effect of stability) smaller.

320

4. Spatial distributions

321 Figure 7 presents spatial distributions of the time-averaged (over the
 322 last hour) surface heat flux obtained from the $a5 \times 5$ case and the $s5 \times 5$
 323 case. Even though both layouts yield similar mean values of the surface
 324 heat flux, it is clear that the staggered case yields a more uniform distri-
 325 bution. Specifically, the difference between the maximum and minimum
 326 surface heat flux is about 0.8% in the staggered wind farm and 2.0% in
 327 the aligned wind farm. In line with experimental measurements (Zhang
 328 et al., 2013a), both layouts show the surface-heat-flux distribution is
 329 heterogeneous. The maximum surface heat flux appears in the region
 330 immediately behind the wind turbine. It should be noted that the
 331 aligned wind farm yields a distinct trend on either side of the column
 332 of turbines, with larger heat flux on the left-hand side of the wind
 333 turbine (looking from the front) and lower heat flux on the right-hand
 334 side of the wind turbine. This is mainly caused by the turbine-induced
 335 flow rotation and the multiple-wake interaction. Flow on the left is on
 336 average moving downward, as also shown previously in Lu and Porté-
 337 Agel (2011) and Porté-Agel et al. (2011), thus transporting high-speed
 338 ($\tilde{u} > \langle \tilde{u} \rangle$) momentum down from higher levels and enhancing mixing
 339 (Fig. 8a). As a consequence, there exists a difference of 0.3 K between
 340 the maximum and minimum temperatures on the land, as shown in
 341 Fig. 8b. Field studies (e.g., Rajewski et al., 2013) have also shown
 342 that the presence of wind turbines can induce heterogeneous spatial
 343 distributions of heat flux and temperature.

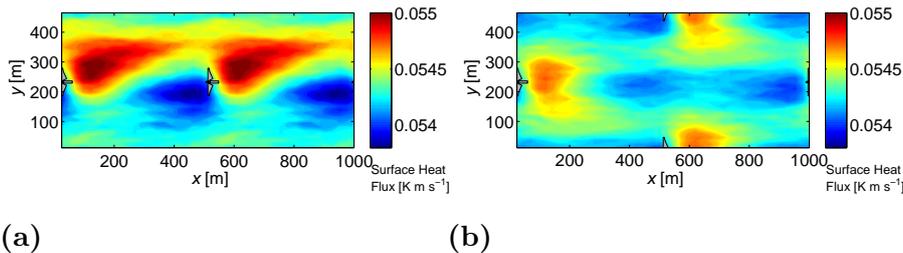


Figure 7. Time-averaged surface heat flux from (a) the $a5 \times 5$ case and (b) the $s5 \times 5$ case.

344 Figures 9 and 10 present the filled contours of the time-averaged
 345 streamwise velocity in a vertical plane and a horizontal plane passing

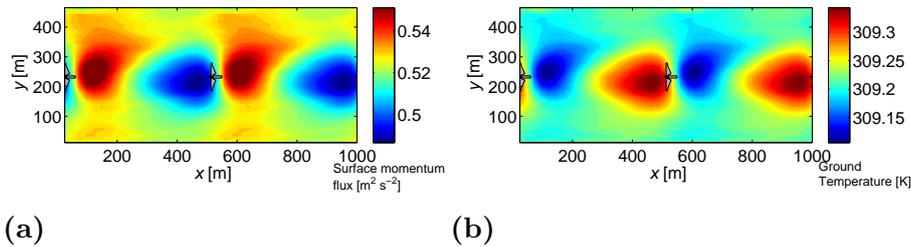


Figure 8. Time-averaged (a) surface momentum flux and (b) ground temperature from the $a5 \times 5$ case.

346 through the turbine axes, obtained from the $a5 \times 5$ case and the $s5 \times 5$
 347 case, respectively. Comparing the two horizontal cross-section contours,
 348 it is clear that the staggered farm yields more uniform flow. Also, in
 349 the staggered farm, the turbine wakes have a longer distance to recover,
 350 which results in a slightly higher wind speed in front of the next down-
 351 stream turbine and, consequently, a larger power output compared with
 the aligned counterpart (shown later).

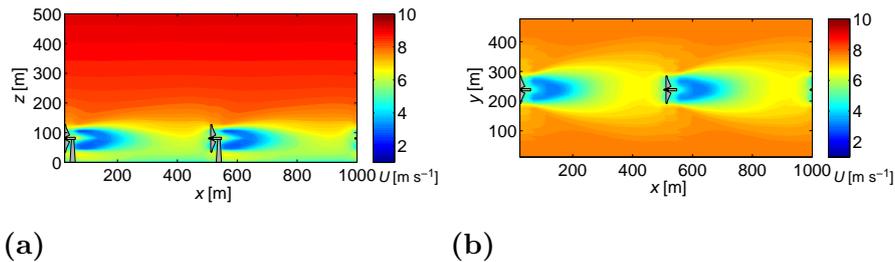


Figure 9. Time-averaged streamwise velocity obtained from the $a5 \times 5$ case: (a) in the vertical plane through the wind-turbine axis; (b) in the horizontal plane at the wind-turbine hub height.

352
 353 The combined effect of velocity deficit/shear and increased turbu-
 354 lence causes increased fatigue loads on wind turbines (e.g., Crespo
 355 et al., 1999; Thomsen and Sørensen, 1999; Vermeer et al., 2003; Eggers
 356 et al., 2003). Figures 11 and 12 show the filled contours of the TKE
 357 in a vertical plane and a horizontal plane passing through the turbine
 358 axes, obtained from the $a5 \times 5$ case and the $s5 \times 5$ case, respectively. In
 359 agreement with previous studies of wind-turbine wakes in boundary-
 360 layer flows (e.g., Wu and Porté-Agel, 2011; Yang et al., 2014; Xie and
 361 Archer, 2014), the turbulence intensity above the hub-height is found
 362 to be larger than that below the hub-height. The primary cause is
 363 the larger velocity shear (and associated production of TKE) observed

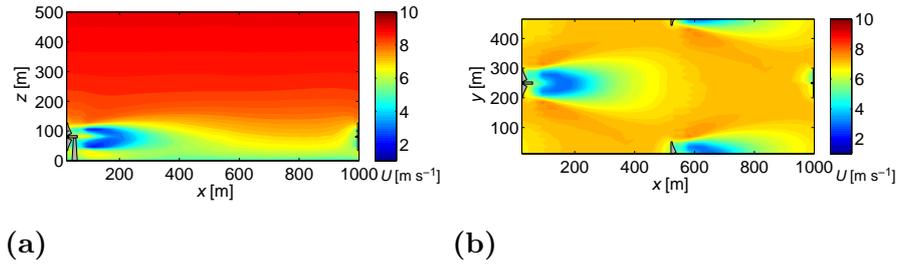


Figure 10. Time-averaged streamwise velocity obtained from the $s5 \times 5$ case: (a) in the vertical plane through the wind-turbine axis; (b) in the horizontal plane at the wind-turbine hub height.

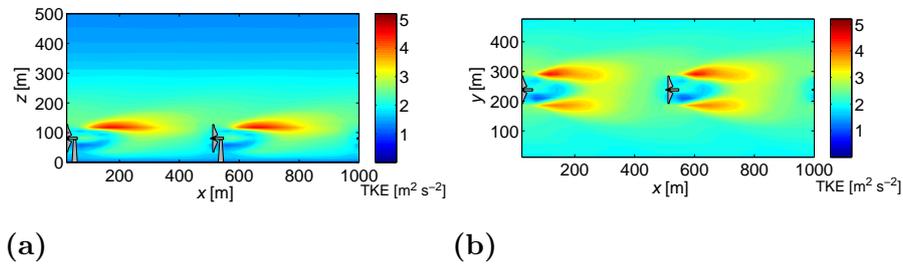


Figure 11. Turbulence kinetic energy obtained from the $a5 \times 5$ case: (a) in the vertical plane through the wind-turbine axis; (b) in the horizontal plane at the wind-turbine hub height.

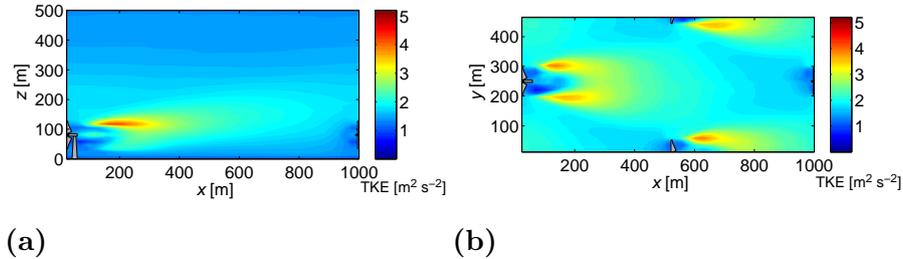


Figure 12. Turbulence kinetic energy obtained from the $s5 \times 5$ case: (a) in the vertical plane through the wind-turbine axis; (b) in the horizontal plane at the wind-turbine hub height.

364 near the turbine top level. The maximum turbulence intensity is found
 365 at that level and at a distance of approximately $1D - 3D$, instead of
 366 $3D - 5D$ observed in stand-alone wind-turbine wakes (Wu and Porté-

367 Agel, 2011; Porté-Agel et al., 2011; Zhang et al., 2013b; Xie and Archer,
 368 2014) and wind-farm wakes in a stable boundary layer (Lu and Porté-
 369 Agel, 2011). The faster recovery of the wind-turbine wakes in the CBL
 370 compared to other situations is also observed in numerical and ex-
 371 perimental studies (e.g., Zhang et al., 2013b; Abkar and Porté-Agel,
 372 2015), and it is attributed to the increased turbulent mixing induced
 373 by convection. Results also show that the magnitude of the maximum
 374 TKE is larger in the aligned layout case ($4.6 \text{ m}^2 \text{ s}^{-2}$), compared with
 375 the staggered one ($3.9 \text{ m}^2 \text{ s}^{-2}$). Like with the velocity deficit, the rela-
 376 tively longer streamwise distance between turbines in the staggered case
 377 allows the turbulence levels to decay to lower values before encountering
 378 the next downwind turbine. It implies that the staggered layout would
 379 yield less fatigue loads on wind turbines. It is interesting to note that,
 380 the maximum added TKE (local TKE maximum behind the turbine
 381 minus TKE level upstream of the turbine) has a similar magnitude
 382 (about $2.0 \text{ m}^2 \text{ s}^{-2}$) for both layouts. It is also similar to the values found
 383 in previous simulations of stand-alone turbine wakes (Abkar and Porté-
 384 Agel, 2015) as well as wind farms in relatively deep neutral boundary
 385 layers (Porté-Agel et al., 2013). It is however much larger than that
 386 reported by Lu and Porté-Agel (2011) for a wind farm in a relatively
 387 shallow stable boundary layer, where the maximum TKE in the wakes
 388 was found to be about $0.6 \text{ m}^2 \text{ s}^{-2}$.

389 5. Temporal measures

390 Figure 13 shows the temporal evolution of the averaged friction velocity,
 391 temperature scale, surface heat flux and heat flux partition into the soil.
 392 The averaging is taken over the horizontal surface plane. Undoubtedly,
 393 wind turbines slow down the wind speed, which in turn yields a decrease
 394 in the friction velocity. Despite the increase in θ_* induced by the wind
 395 farms, the net effect is a reduction of 4%-7% in the surface heat flux.
 396 The reduction is slightly stronger with increasing wind-turbine density.
 397 As a result of the reduction in the surface heat flux, more energy is
 398 available to heat up the soil, causing higher land-surface temperature
 399 as shown in Table II.

400 Figure 14 shows the temporal evolution of the boundary-layer height,
 401 the averaged convective velocity scale ($w_* = \left(\frac{g}{\Theta_0} u_* \theta_* z_i\right)^{\frac{1}{3}}$), large-eddy
 402 turnover time (z_i/w_*) and stability parameter ($-z_i/L$). The effects of
 403 wind-turbine spacing and layout on the first three variables are not
 404 significant. All cases show approximately 16% growth of the boundary
 405 layer, and 10% increase in updraft speed and turn-over time. The

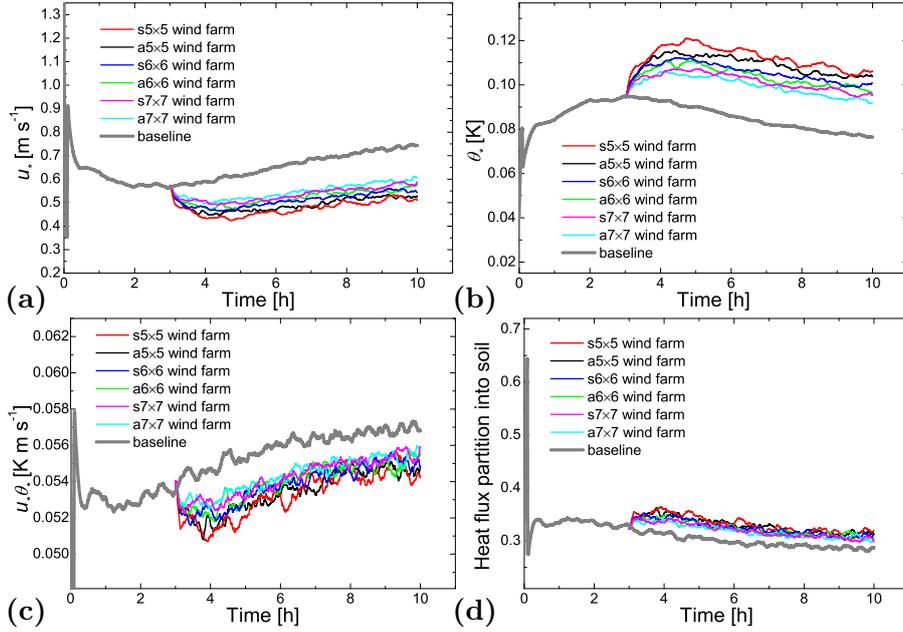


Figure 13. Evolutions of (a) friction velocity, (b) temperature scale, (c) surface heat flux, and (d) heat flux partition into soil.

406 increased value of the stability parameter indicates that the wind-farm
 407 wake significantly changes the static stability of the flow, as shown
 408 above in Fig. 6. Not only the boundary-layer height is increased, but
 409 also the magnitude of the Obukhov length is significantly reduced. This
 410 is because the Obukhov length is computed using surface fluxes, and the
 411 reduction in the magnitude of the surface momentum flux is relatively
 412 stronger compared with the reduction in the surface heat flux.

413 For each wind turbine, its power production can be related to the
 414 incoming wind speed measured at hub height (Bozkurt et al., 2014).
 415 Here, we adopt the SWT-2.3-93 power curve from the manufacturer
 416 to compute the wind-turbine power output using the wind speed $1D$
 417 upwind of the turbine. Figure 15 shows the temporal evolution of the
 418 wind-power production, which is averaged over all the wind turbines.
 419 We summarize the mean normalized wind-turbine power output in Ta-
 420 ble III . It is clear that, as expected, the lower the turbine density, the
 421 higher the power extracted by each individual turbine. Moreover, for
 422 the same turbine density, results show that the staggered wind farms
 423 yield about 10% more power output than their aligned counterparts. A
 424 similar increase has also been reported by Abkar and Porté-Agel (2013)
 425 in simulations of a conventionally-neutral ABL.

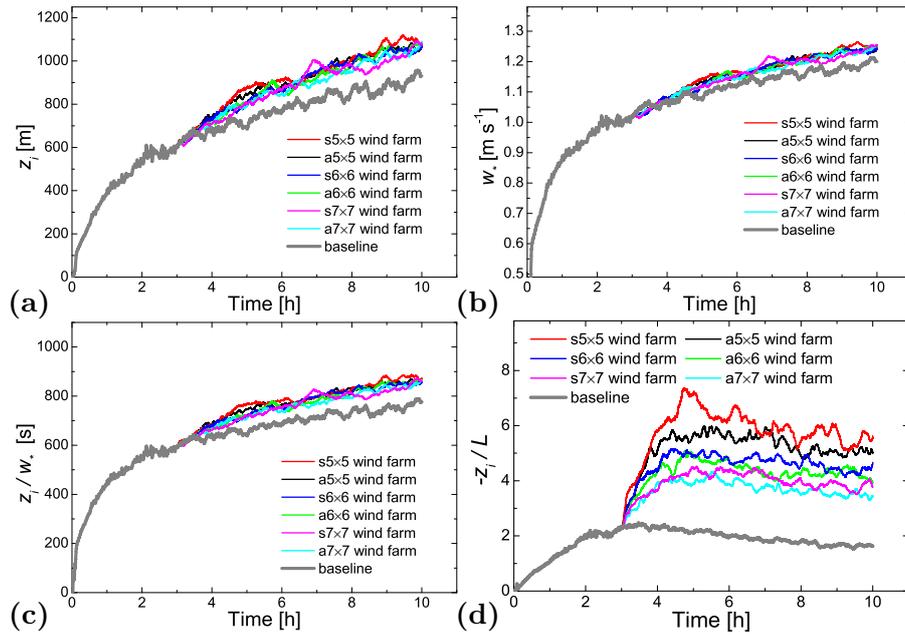


Figure 14. Evolutions of (a) boundary-layer height, (b) convective velocity scale, (c) large-eddy turnover time, and (d) stability parameter.

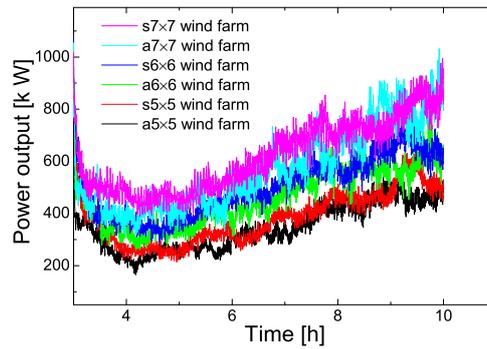


Figure 15. Averaged wind-turbine power output.

Table III. Mean normalized (by mean wind-turbine power output from the a5x5 wind farm) wind-turbine power output.

Case	a5x5	s5x5	a6x6	s6x6	a7x7	s7x7
Normalized Power	100%	110.8%	131.6%	143.4%	162.6%	176.4%

6. Conclusions

427 In this study, large-eddy simulations have been used, for the first time,
428 to investigate the effects of very large wind farms on a CBL. The sim-
429 ulation results show that the wind-turbine wakes enhance the vertical
430 mixing, resulting in changes to the ABL flow stability. Results also
431 reveal that wind farms lead to a slight reduction in the surface heat
432 flux, an increase in land-surface temperature, and a slight reduction in
433 the vertically-integrated temperature. These effects increase with in-
434 creasing wind-turbine density. It should be noted here that the warming
435 effect on the land surface is due to the redistribution of energy available
436 for the soil and the air; in particular, a reduction in the sensible heat
437 flux is associated with an increase in the heat flux into the soil. Results
438 also show that the spatial distribution of the surface heat flux is het-
439 erogeneous. In addition, the connection between the surface heat-flux
440 heterogeneity and the coherent wake column rotation in aligned wind
441 farms suggests that it is essential to simulate the wake rotation effects
442 in numerical models of turbine-wake flows in order to reproduce the
443 spatial distribution of the surface heat flux.

444 The staggered wind-farm layout is characterized by a relatively longer
445 separation between consecutive downwind turbines compared to the
446 aligned counterpart. As a result, the staggered layout allows the wakes
447 to recover more, exposing the downwind turbines to higher local wind
448 speeds (leading to higher wind-power production) and lower turbulence
449 intensity levels (leading to lower fatigue loads).

450 This study provides evidence and quantification of the impact of
451 wind farms on daytime convective boundary layers. This information is
452 of great importance for optimizing the design of wind farms and also for
453 developing improved parametrisation of turbulent fluxes in weather and
454 climate models. Future work will focus on further investigating wind-
455 farm-atmosphere interactions under more realistic conditions, including
456 complex terrain (e.g. topography and/or surface heterogeneity), non-
457 stationary flow conditions, and finite-size wind farms.

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